

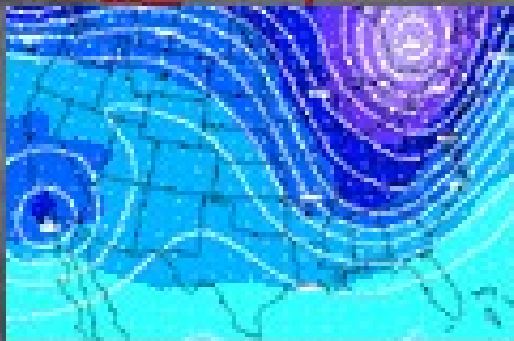
DEFINITION OF A CYCLONE

Midlatitude Cyclones

WINDS
ASSOCIATES WITH
A CYCLONE



UPPER AIR FEATURES



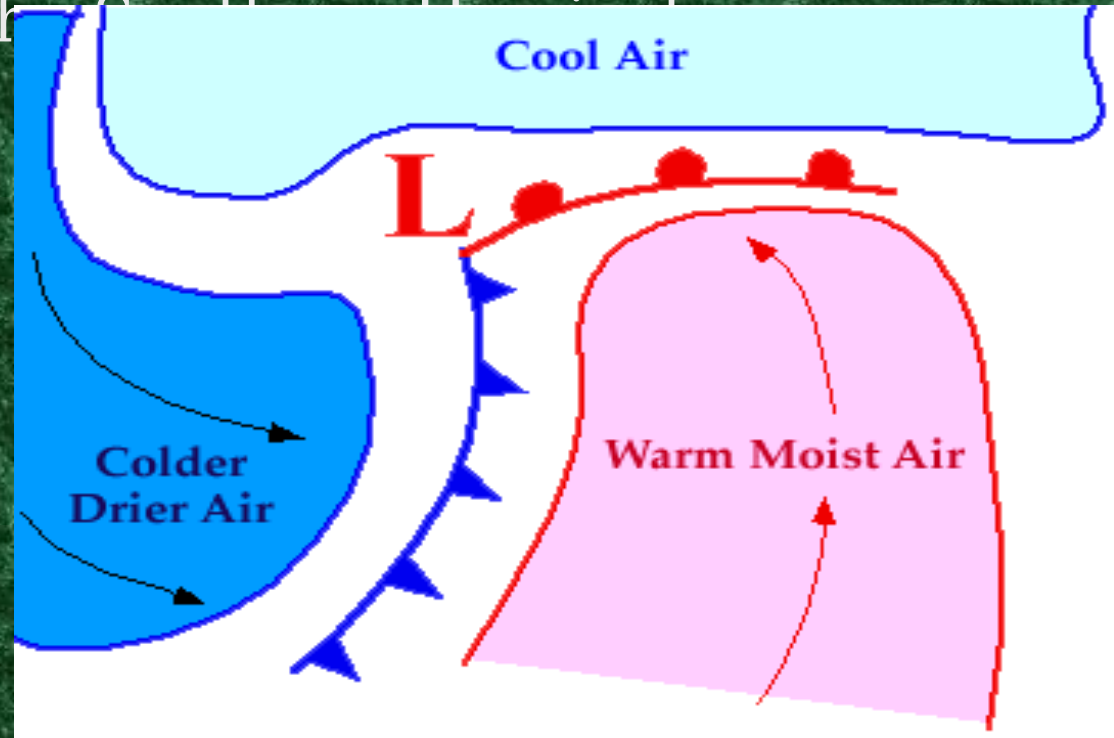
AIR MASSES &
CYCLONES

CYCLONES ON
SATELLITE IMAGES

Cyclones

an idealized
model

A cyclone is an area of **low pressure** around which the winds flow counterclockwise in the Northern Hemisphere and clockwise in the Southern Hemisphere.

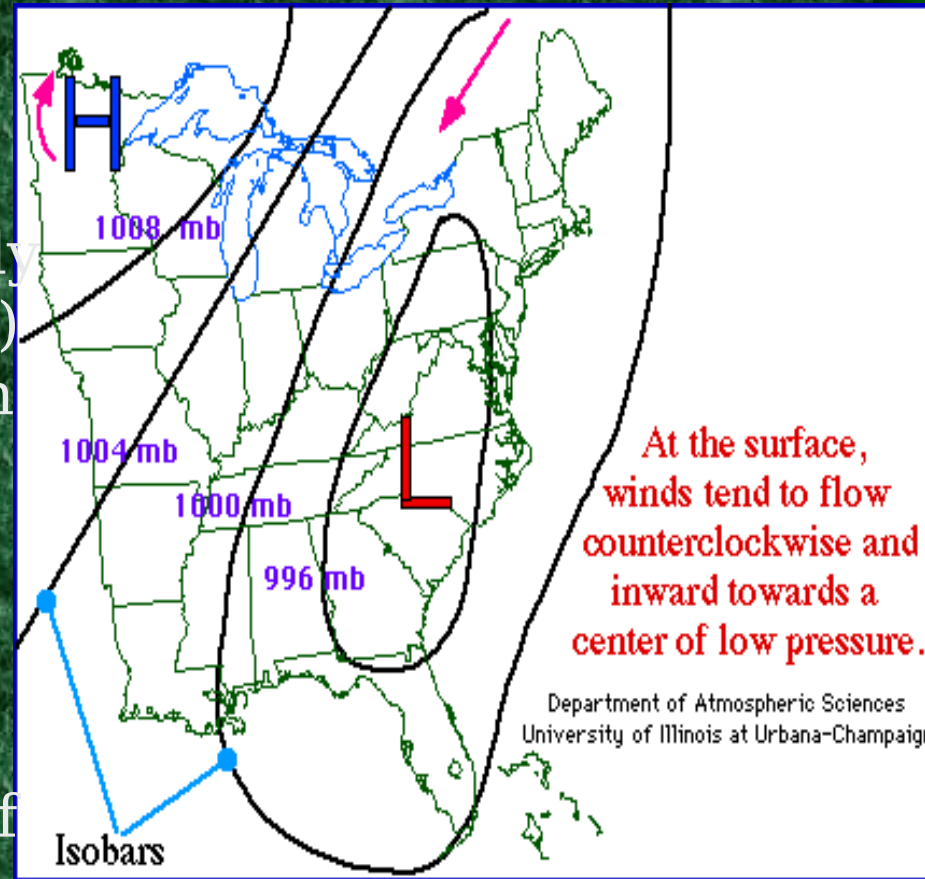


A developing cyclone is typically accompanied by a **warm front** pushing northward and a **cold front** pulling southward, marking the leading edges of air masses being wrapped around a center of **low pressure**, or the

The counterclockwise winds associated with northern hemisphere midlatitude cyclones play a significant role in the movement **air masses**, transporting warm moist air northward ahead of a low while dragging colder, drier air southward behind it.

Rising air in the vicinity of a low pressure center favors the development of clouds and precipitation, which is why cloudy weather (and likely precipitation) are commonly associated with an area of low pressure. Cyclones are easily identifiable on certain types of weather maps by remembering some key signatures. For example, a cyclone can be found on a map of surface observations by recognizing a

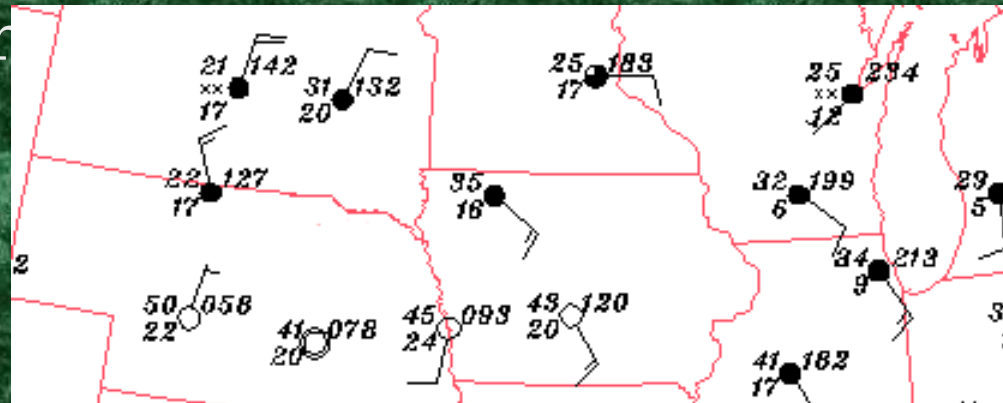
counterclockwise rotation of the



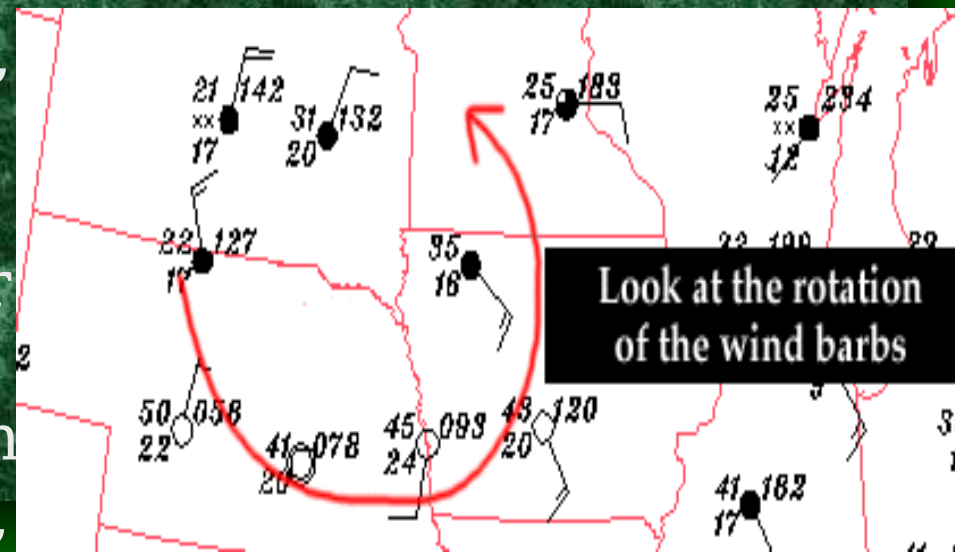
Winds Around Cyclones

flowing counterclockwise in the northern
hemisphere

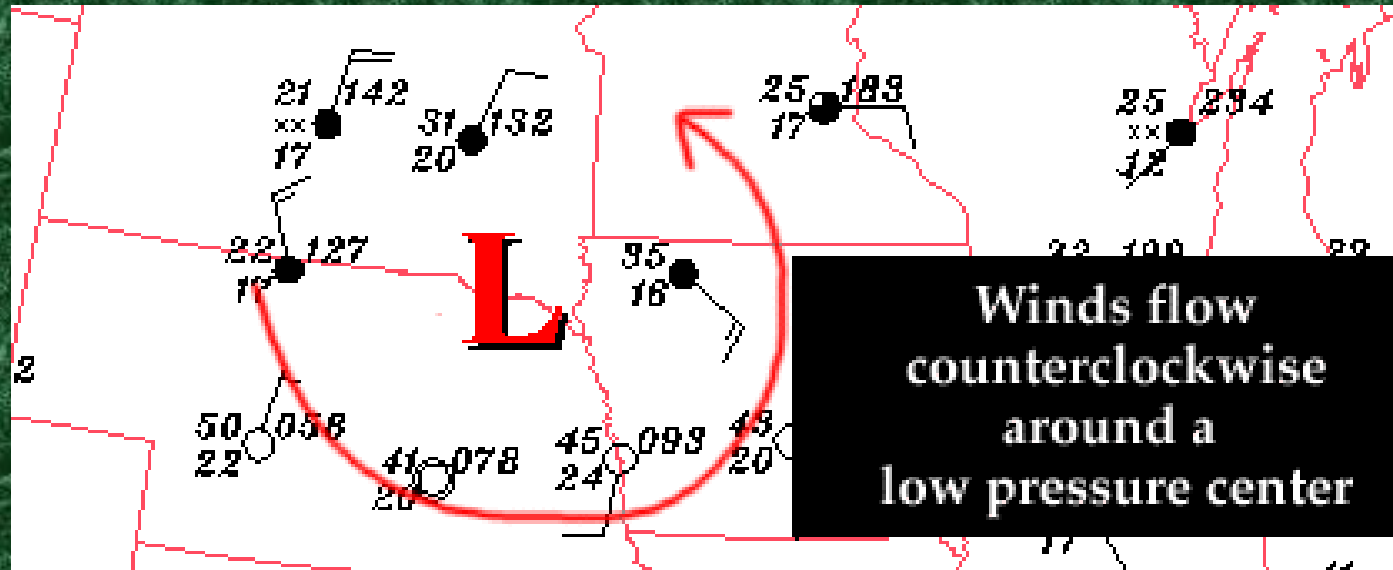
Wind barbs are useful for locating **low pressure centers**
on surface weather



Since winds flow in a
counterclockwise direction
around **low pressure centers**,
look for a group of stations
where the **wind barbs** reflect
this type of wind pattern. For
example, a counterclockwise
wind pattern was observed in
the states of Nebraska, Iowa,



The **low pressure center** was located near the center (similar to the center of a whirlpool) with winds flowing counterclockwise around it.



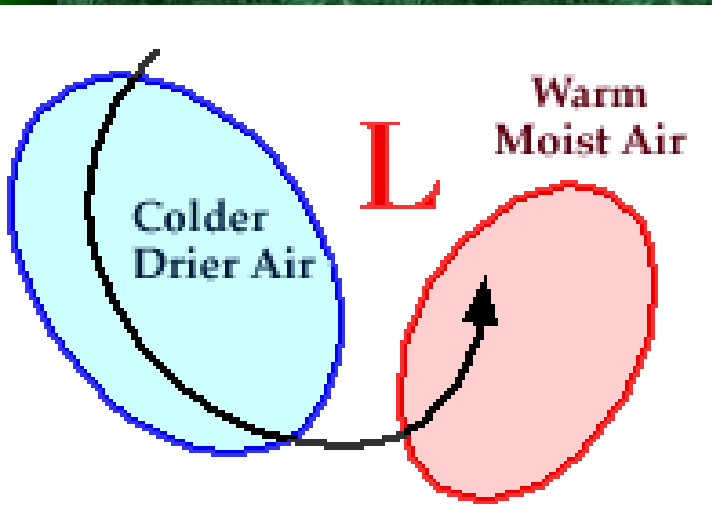
So when trying to find a **low pressure center** on a surface weather map, use the **wind barbs** to identify a counterclockwise wind pattern and the low pressure center will be found near the center of circulation.

The Movement of Air

Masses

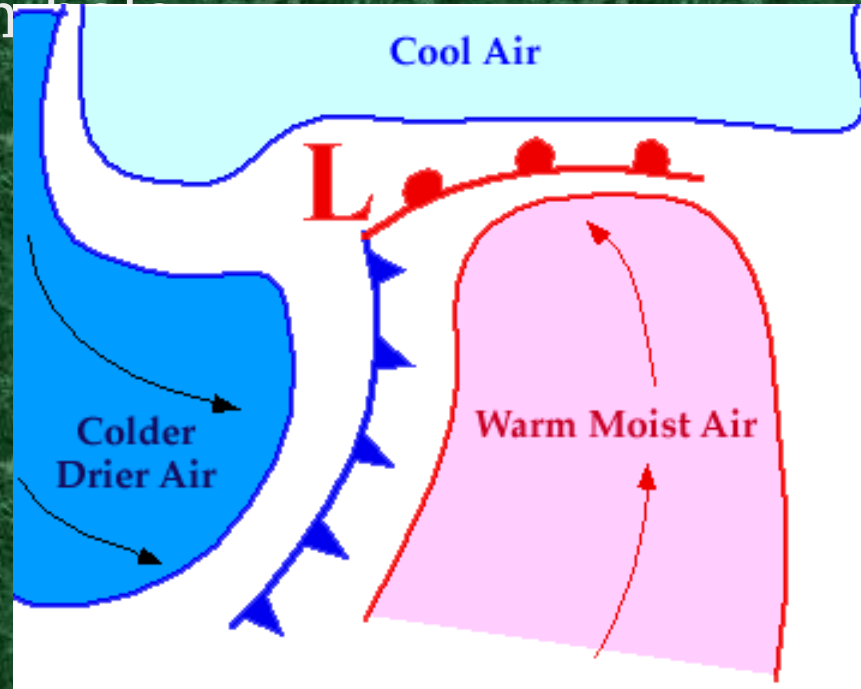
transporting warm air northward and colder air southward

Counterclockwise winds associated with cyclones transport heat and moisture from lower to higher latitudes and play a significant role in the movement of air masses.



As a cyclone intensifies, (the central pressure drops), counterclockwise winds around the low pressure center also intensify, transporting the air masses around the center of circulation.

By superimposing **fronts** over the **low pressure center** and the **air masses**, a top view of a midlatitude cyclone and accompanying air masses might resemble something like the diagram below.



Southerly winds east of the low transport warm and moist air northward and this moisture often contributes to the development of precipitation. A **warm front** marks the leading edge of this warm, moist air mass. Behind the low, northerly winds transport colder and drier air southward, with a **cold front** marking the leading edge of

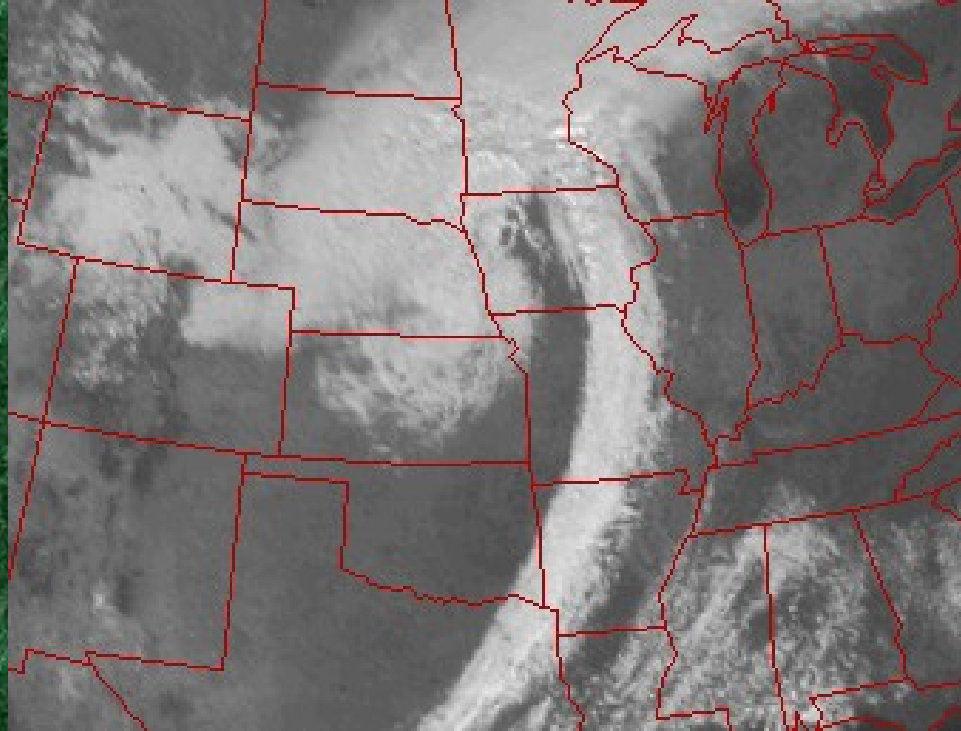
Cyclones on Satellite

Images

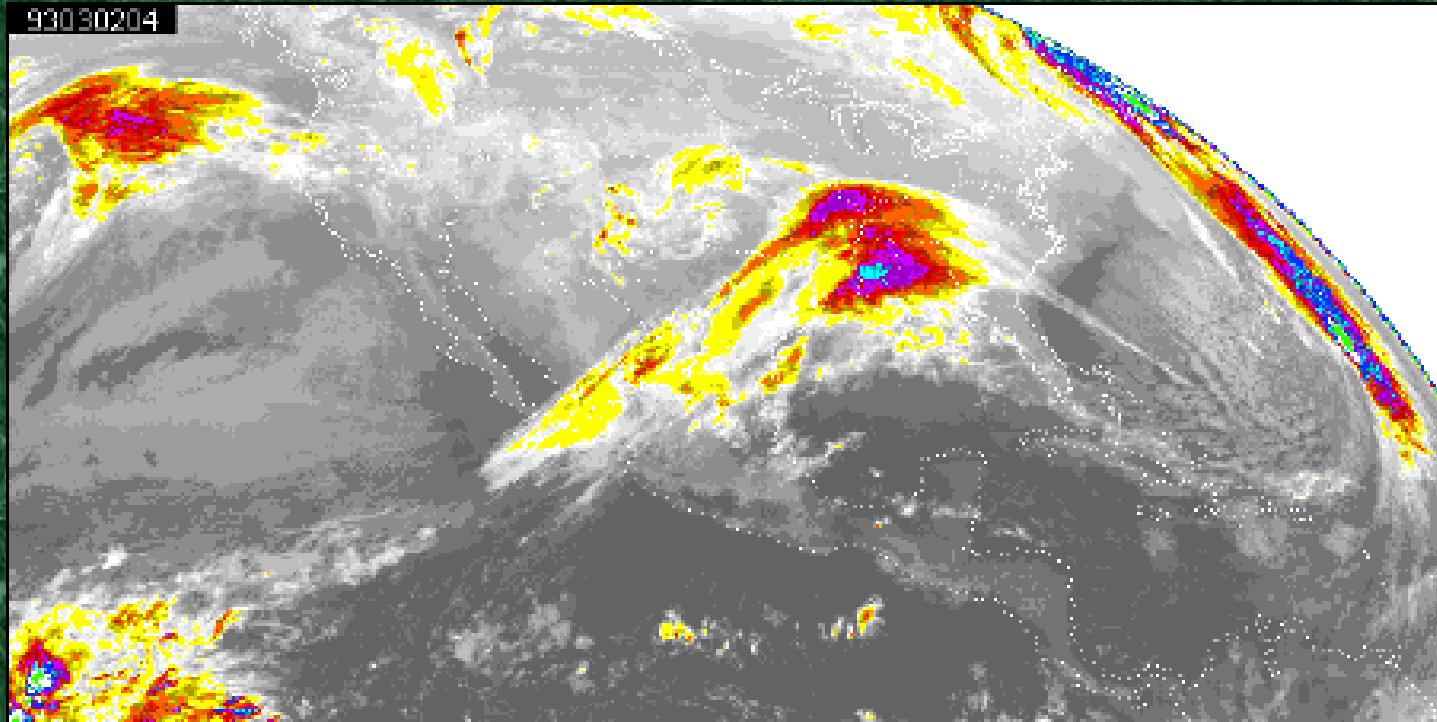
comma-shaped cloud

configuration

On satellite images, a **midlatitude** cyclone is often identifiable by a comma-shaped cloud mass



A single **cyclone** can influence the weather over a large area, (in this case from Texas into Minnesota). This particular storm (in the satellite image above) left more than six inches of snow from Nebraska into Minnesota, while heavy rains occurred from Missouri into Texas.



The center of the cyclone was initially located over the southcentral portions of the United States. As the cyclone evolved over time, notice the counterclockwise rotation of clouds around the **cyclone center**.

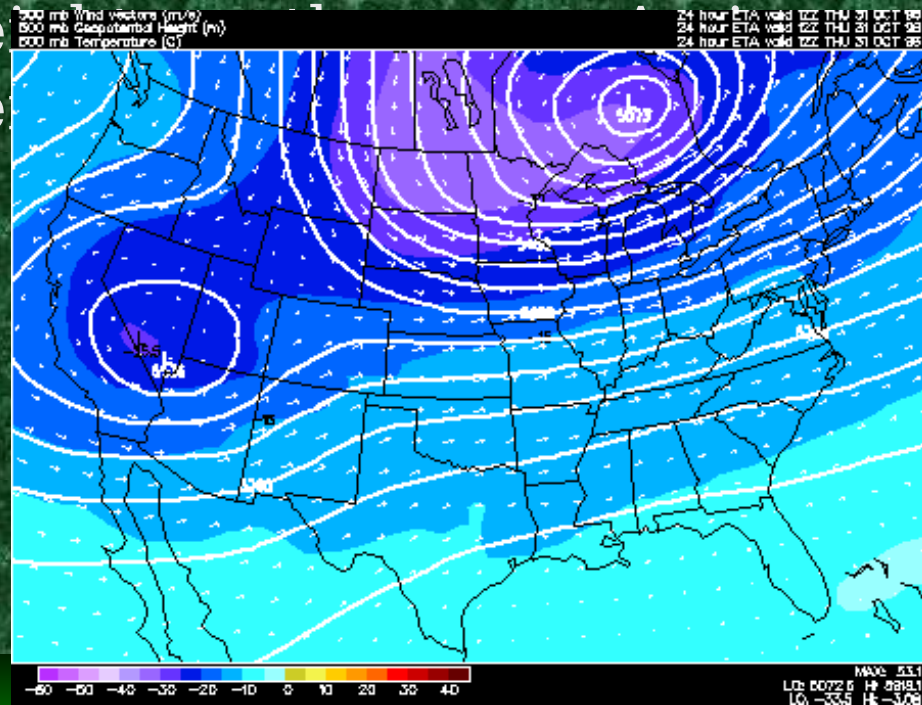
Geopotential

Height

height of a given

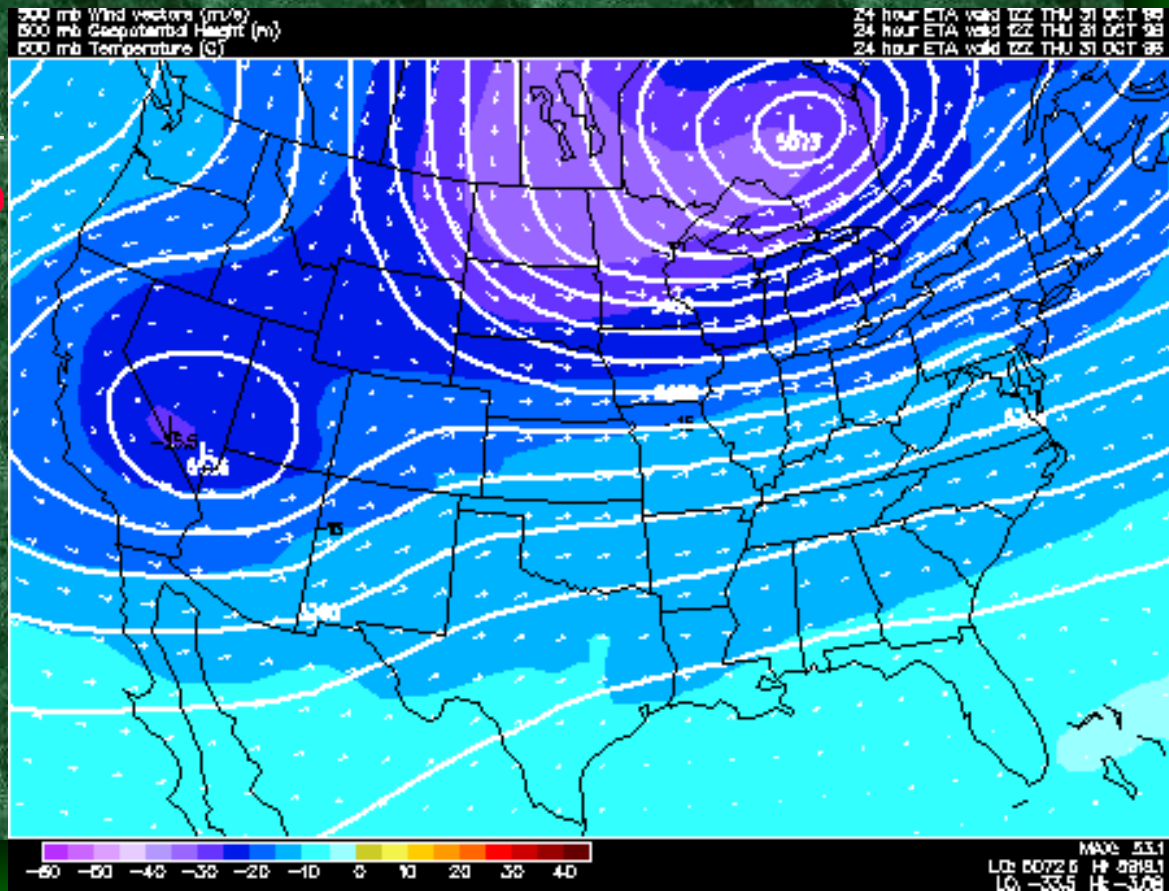
pressure

Geopotential height approximates the actual height of a pressure surface above mean sea-level. Therefore, a geopotential height observation represents the height of the pressure surface on which the observation was taken. A line drawn on a weather map connecting points of equal height (in meters) is called a height contour. That means, at every point along a given contour, the values of geopotential height are the same. The following map is depicting the



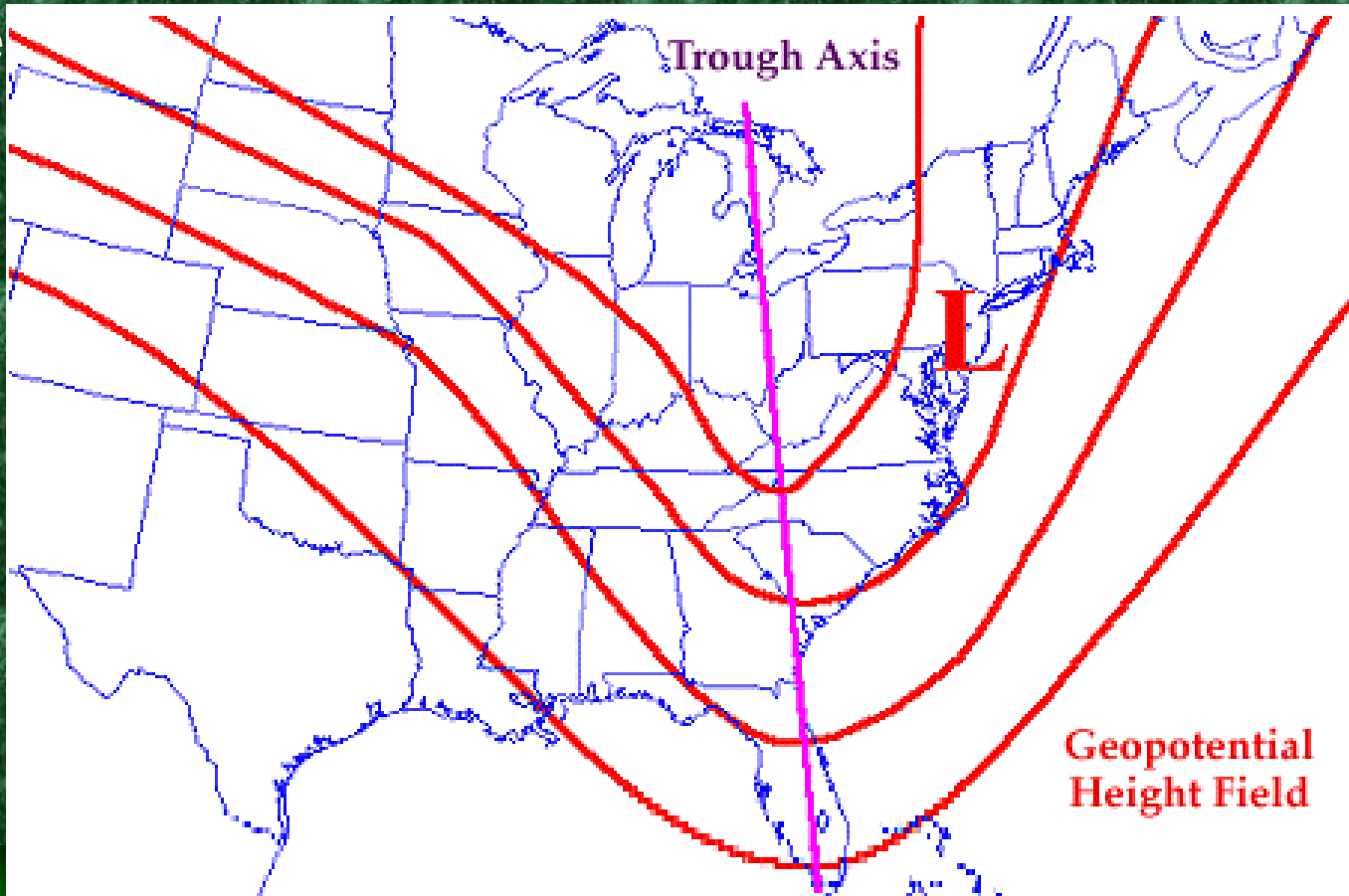
Height contours are represented by the solid lines. The small numbers along the contours are labels which identify the value of a particular height contour (for example 5640 meters, 5580 meters, etc.). This example depicts the 500 mb geopotential height field and temperatures (color filled regions). The height field is given in meters with an interval of 60 meters.

Geopotential height is **ridges** which are the **u** **cyclones** and **anticyclo**



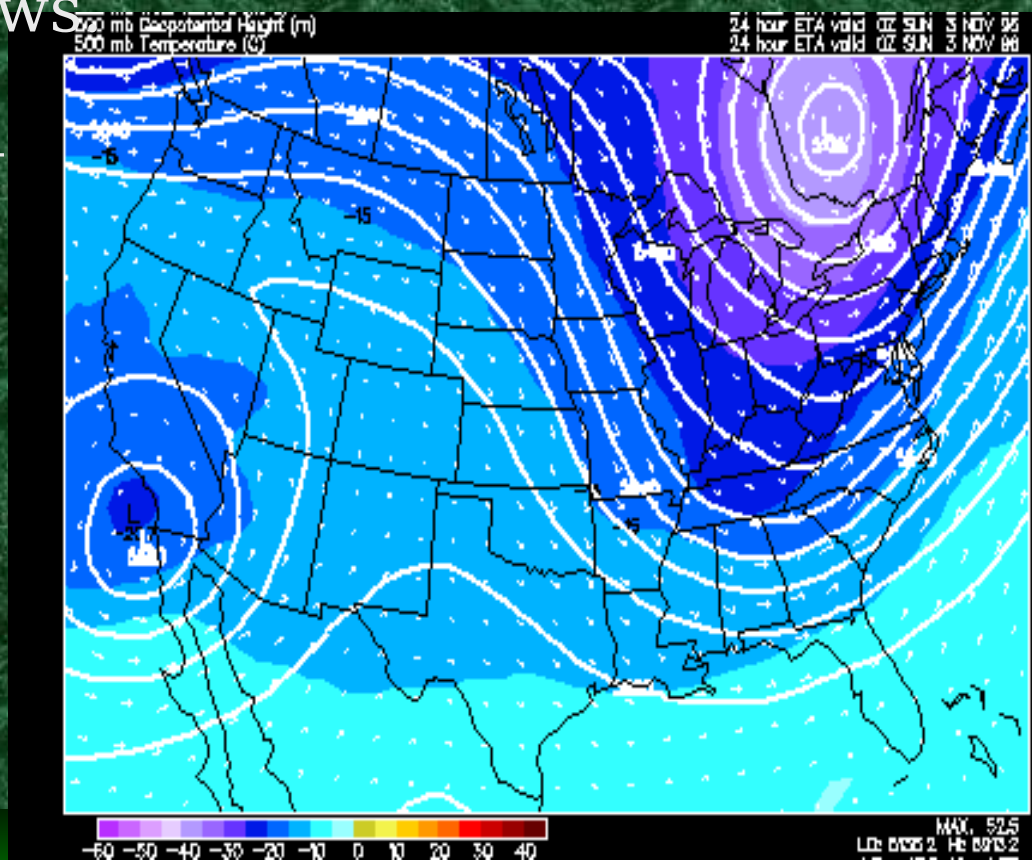
Trough
upper level
lows

When the **height contours** bend strongly to the south, (as in the diagram below), it is called a TROUGH. Strong troughs are typically preceded by stormy weather and colder air at the surface. Below is an example of a trough in an upper-level height field (red contours). The trough axis is de



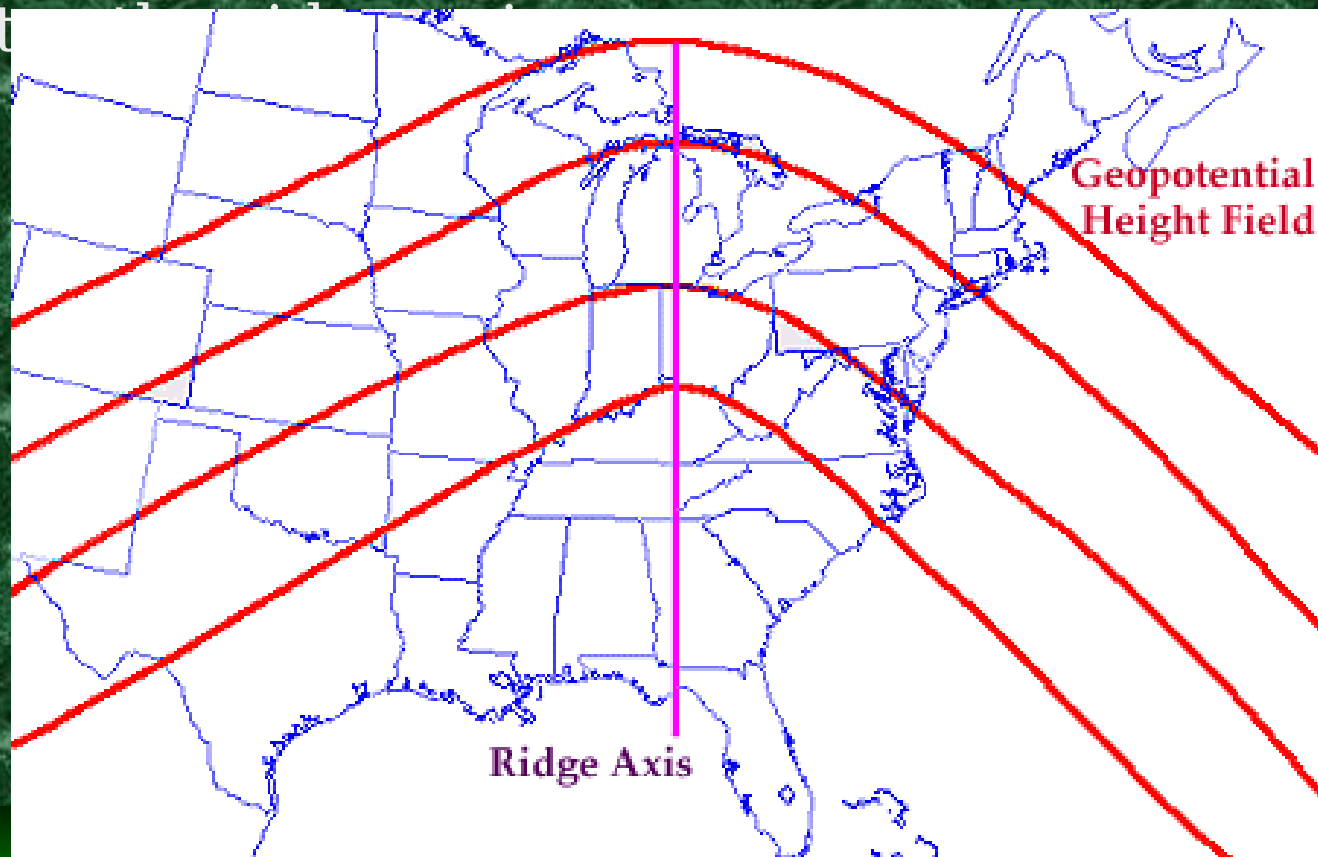
The image below depicts geopotential height (solid white contours) and temperatures (colored regions) at 500 mb. Temperatures decrease with color from light blue to purple. A trough is located over the eastern United States and is indicated by the dip in the **geopotential** height field. This is the upper level extension of a surface **low pressure center**, which is why troughs are also called upper level lows.

Notice the relatively cold temperatures associated with the trough. This is caused by the southward transport of colder air in the lower troposphere. The trough will intensify (deepen further southward) if cold air continues to move



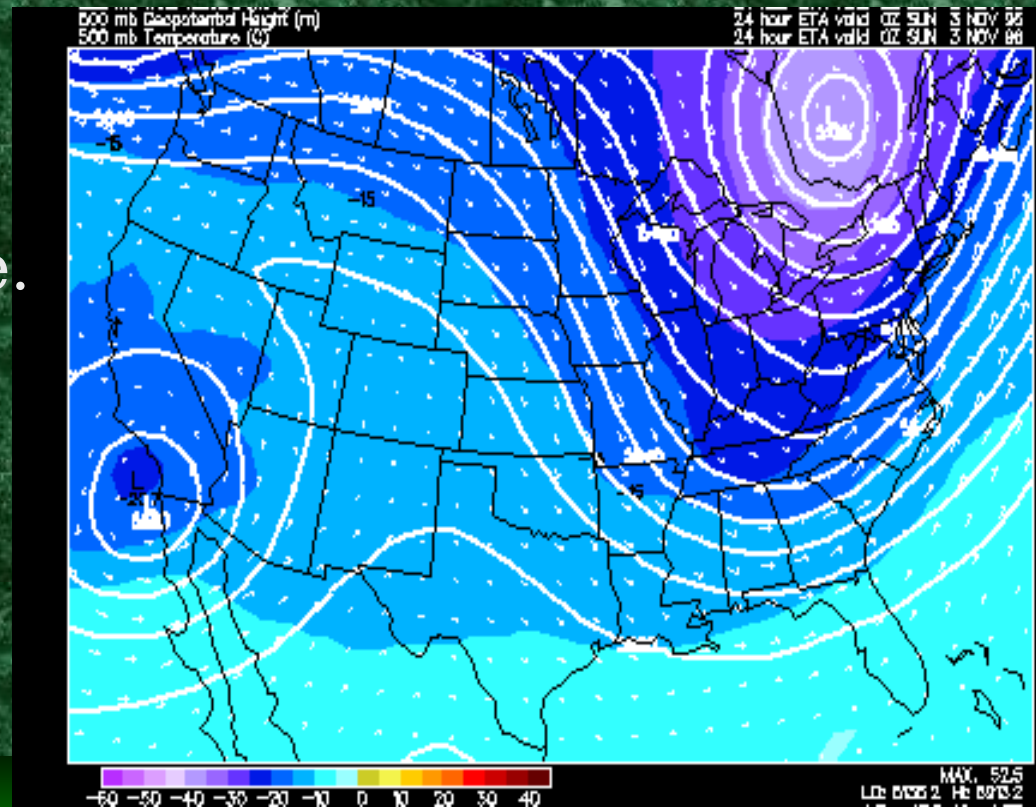
Ridges upper level highs

When the **height contours** bend strongly to the north (as in the diagram below), this is known as a RIDGE. Strong ridges are accompanied by warm and dry weather conditions at the surface. Below is an example of a ridge in an upper-level height field (red contours). The purple line denotes the ridge axis.



The image below depicts geopotential height (solid white contours) and temperatures (colored regions) at 500 mb. Temperatures decrease with color from light blue to purple. A ridge is located from Texas into Montana and is indicated by the bulge in the **geopotential** height field. This is the upper level extension of a surface **high pressure center**, which is why ridges are also called upper level highs.

Notice the relatively warm temperatures associated with the ridge. This is caused by the northward transport of warmer air in the lower troposphere. The ridge will intensify (bulge further northward) if warm air continues to



Trough and Ridge

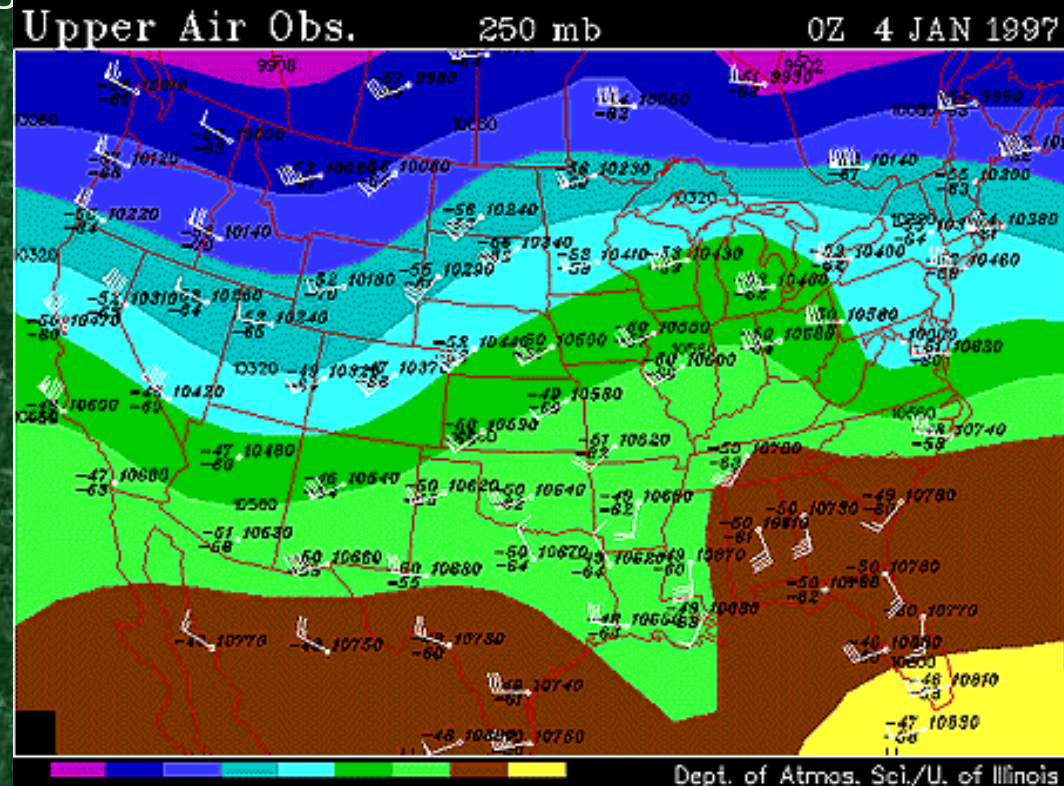
Amplification

in response to lower level

temperature advection

Warm advection beneath an upper level ridge causes it to build (increase in amplitude), while **cold advection** beneath an upper level **trough** will contribute to its deepening. The animation below highlights the amplification of ridge/trough system at 250 mb in response to intense **thermal advection** at 850 mb. The total time elaps

Observe how the **ridge** over the east gradually builds (raises northward) while the **trough** over the northern plains gradually deepens (sinks southward). While the entire ridge/trough system propagates eastward. Such amplification of upper level waves can increase the



Rising Motion and Surface Pressure Falls

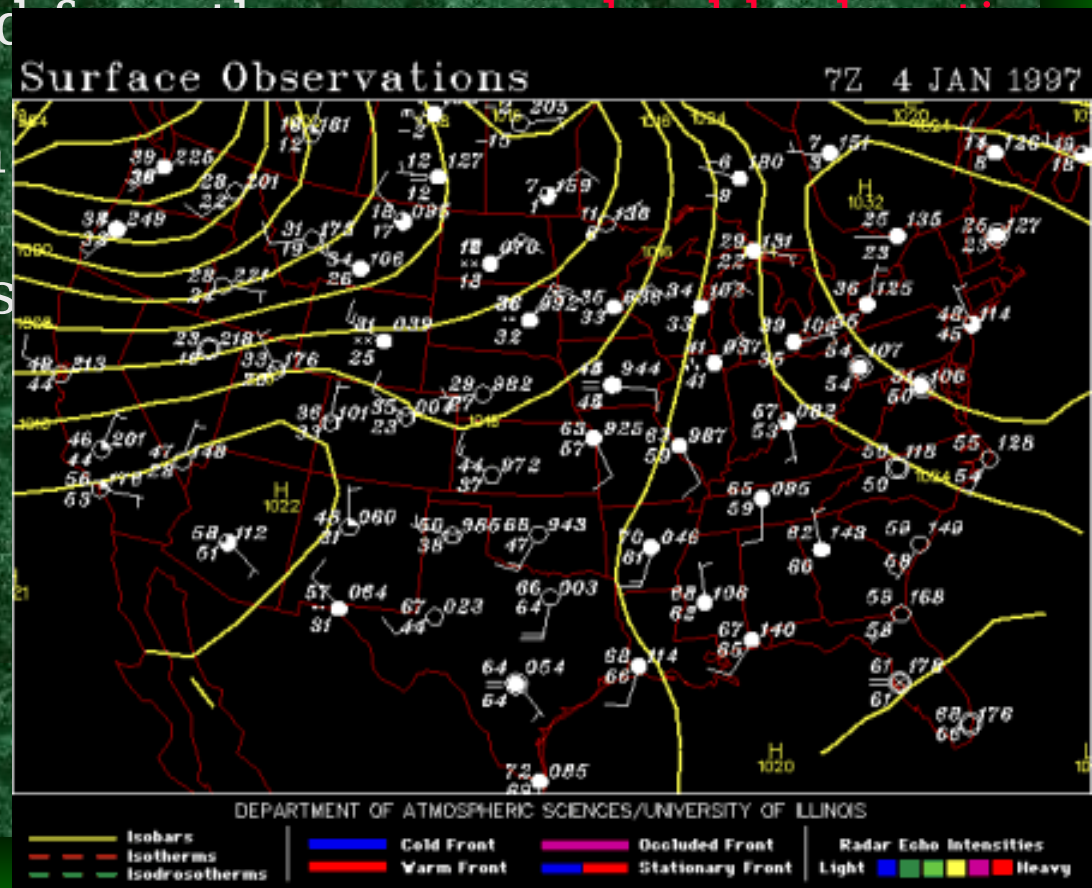
in response to warm and cold

advections

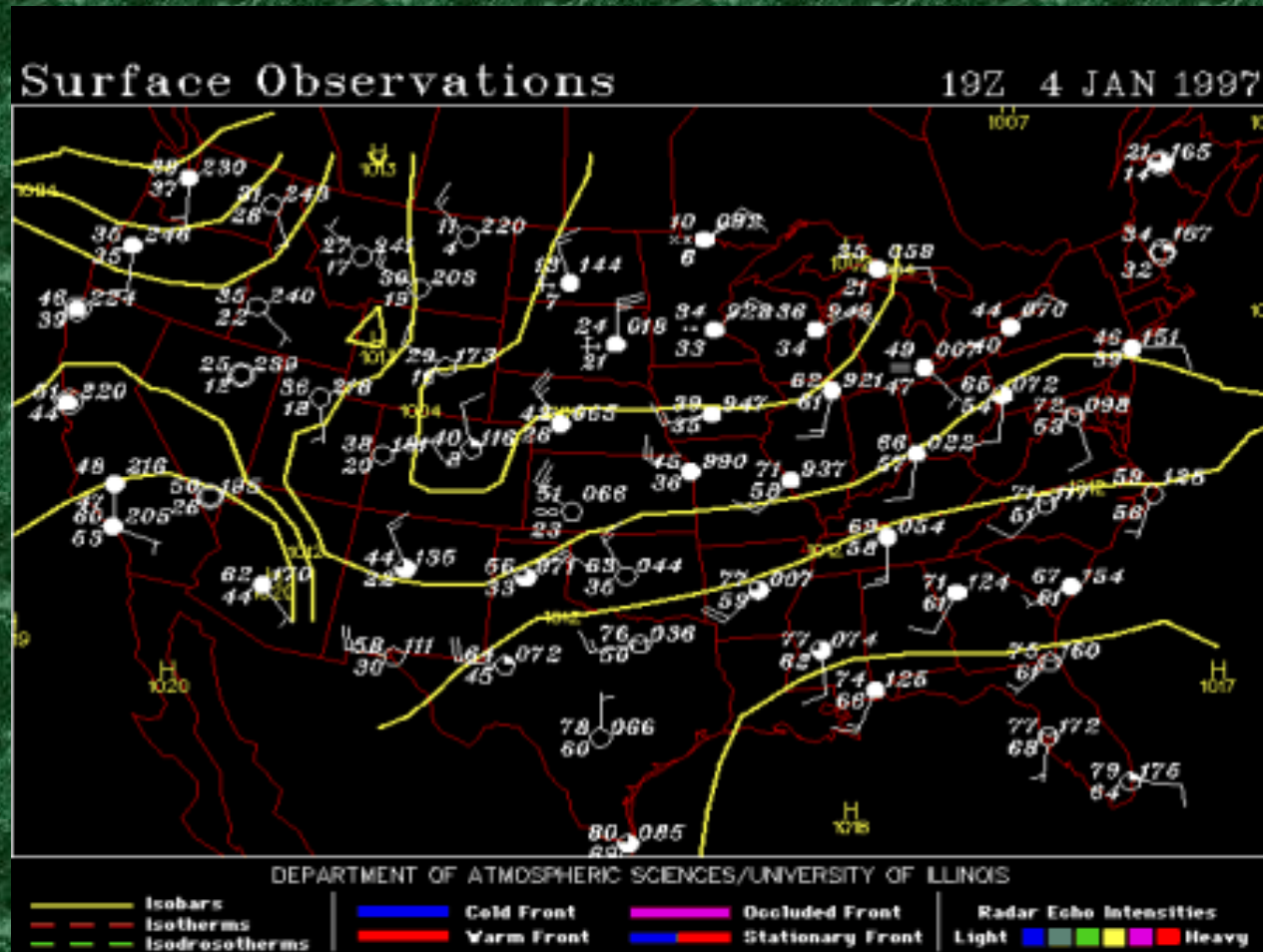
Warm and cold advection influence vertical air motion.

Warm advection results in rising motion which leads to falling pressures at the surface, while cold advection leads to sinking motion, causing pressures to rise at the surface. The sequence of surface maps below show the surface pressure fields (isobars) that resulted from this process.

The top map shows an area of low pressure over the northwestern United States and an area of high pressure over the eastern U.S. 12 hours later, there is a noticeable decrease in pressure from Texas to Illinois (region of strongest warm advection at 850 mb) while surface



A **cyclone** at the surface that moves under an area of warm advection at 850mb is likely to deepen. For this reason, systems at the surface will tend to "phase lock" with systems aloft, and they will propagate more or less together.

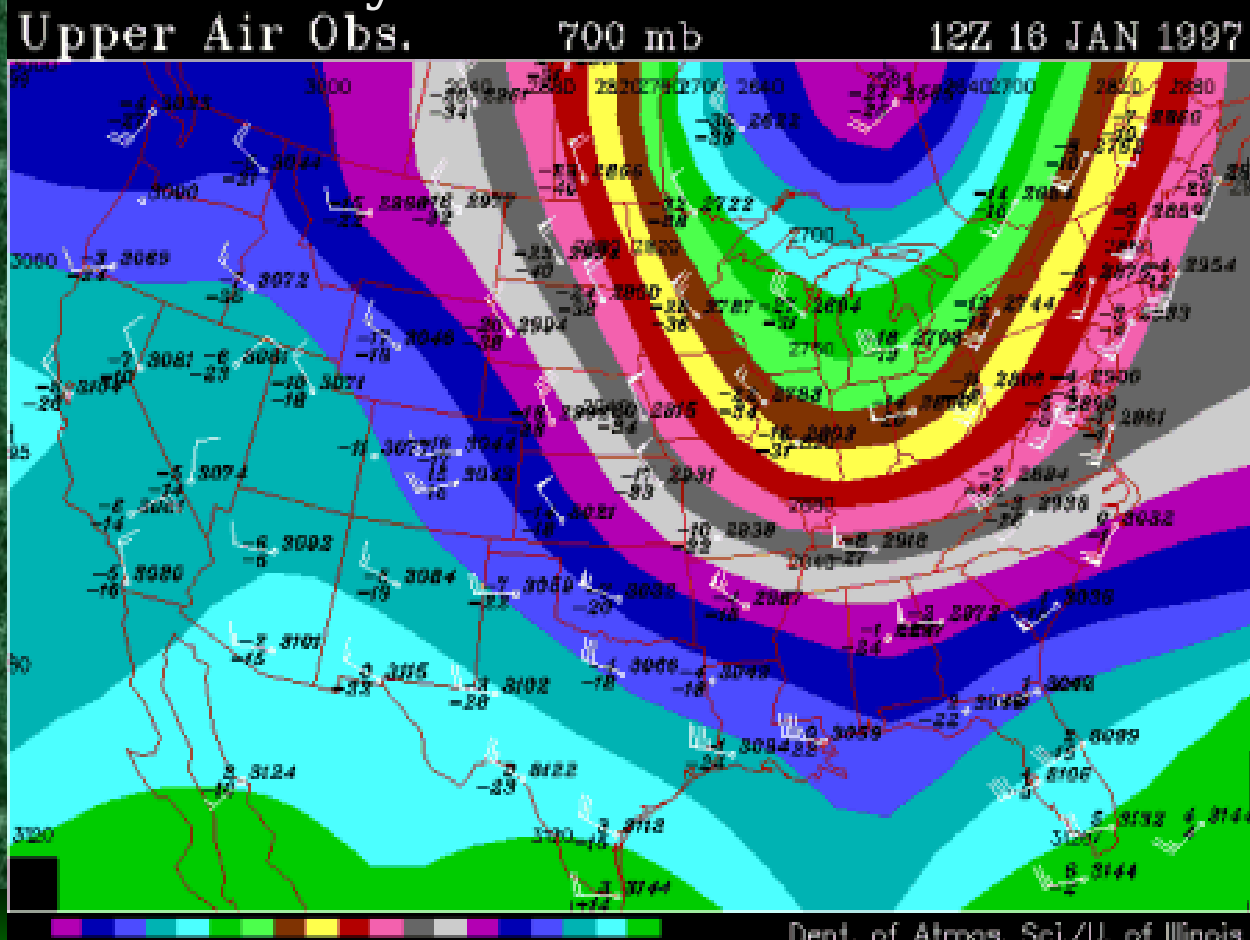


Steering Level For Cyclones

guided by winds near
700 mb

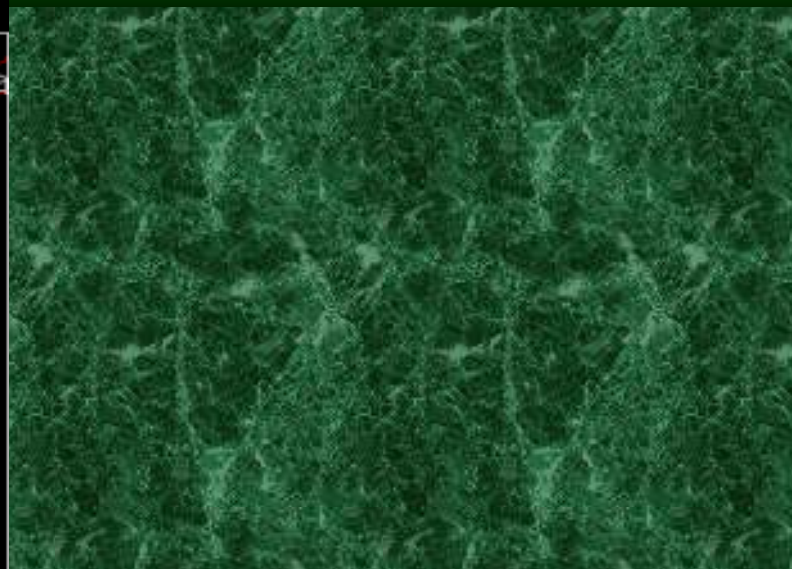
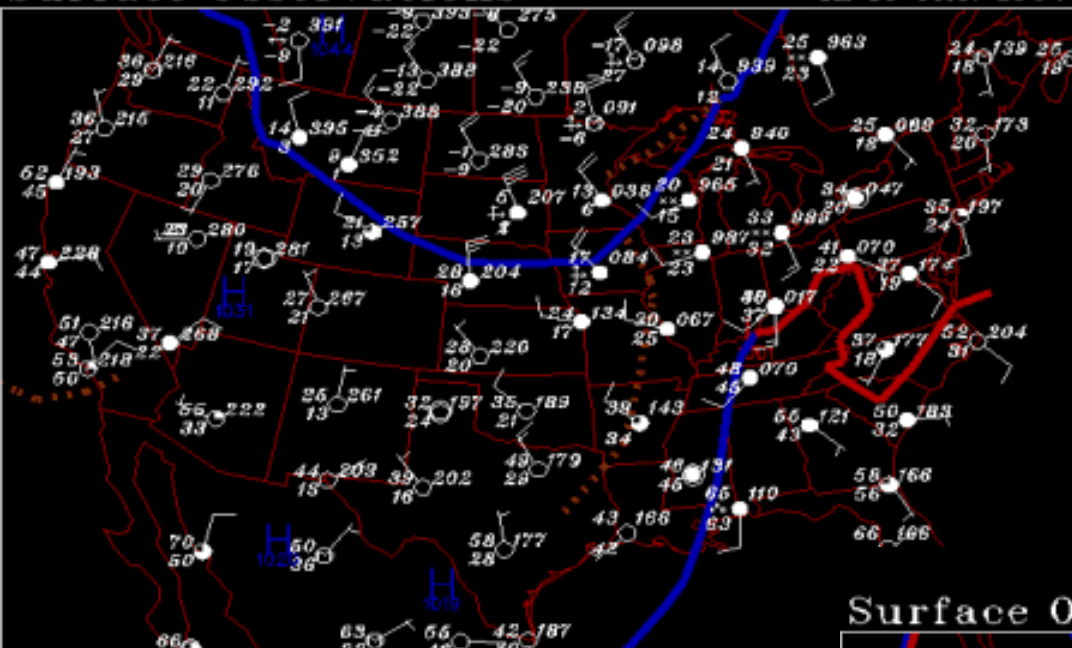
Surface systems often follow the prevailing airflow at this level. Therefore, the airflow at 700mb, inferred from the **geopotential** contours and geostrophic balance, can indicate the direction surface systems will take.

Notice that winds over the eastern U.S. are generally out of the south. At the same time, a **surface cyclone** was centered over the Ohio Valley.



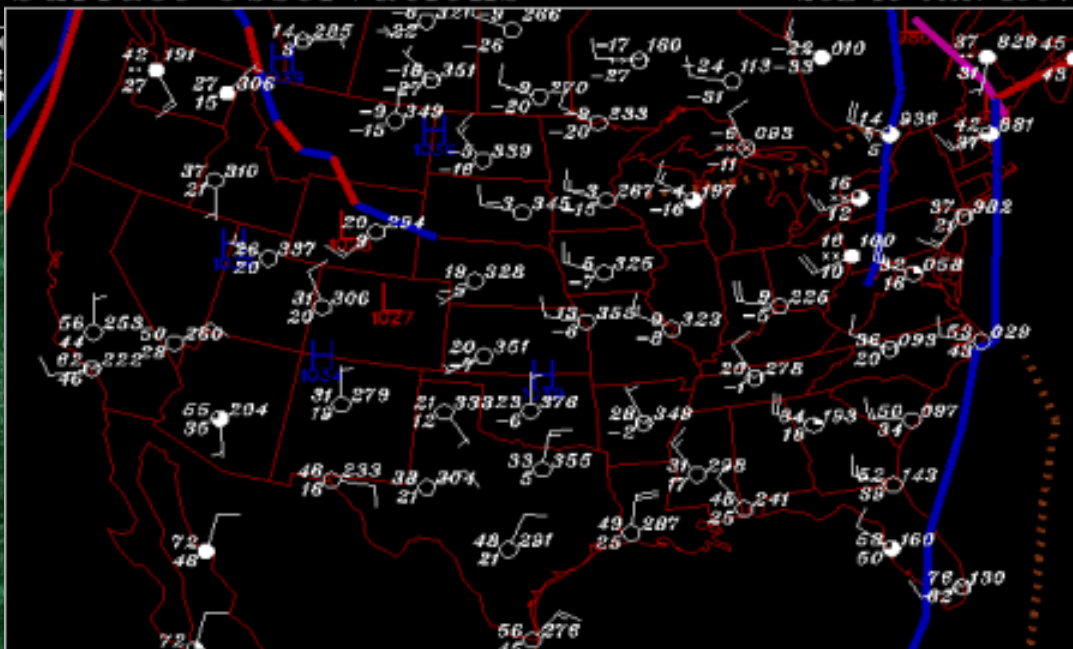
Surface Observations

12Z 16 JAN 1997



Surface Observations

23Z 16 JAN 1997



24 hours later, the cyclone propagated northeastward (parallel to the winds at 700mb) and was centered over Eastern Canada.



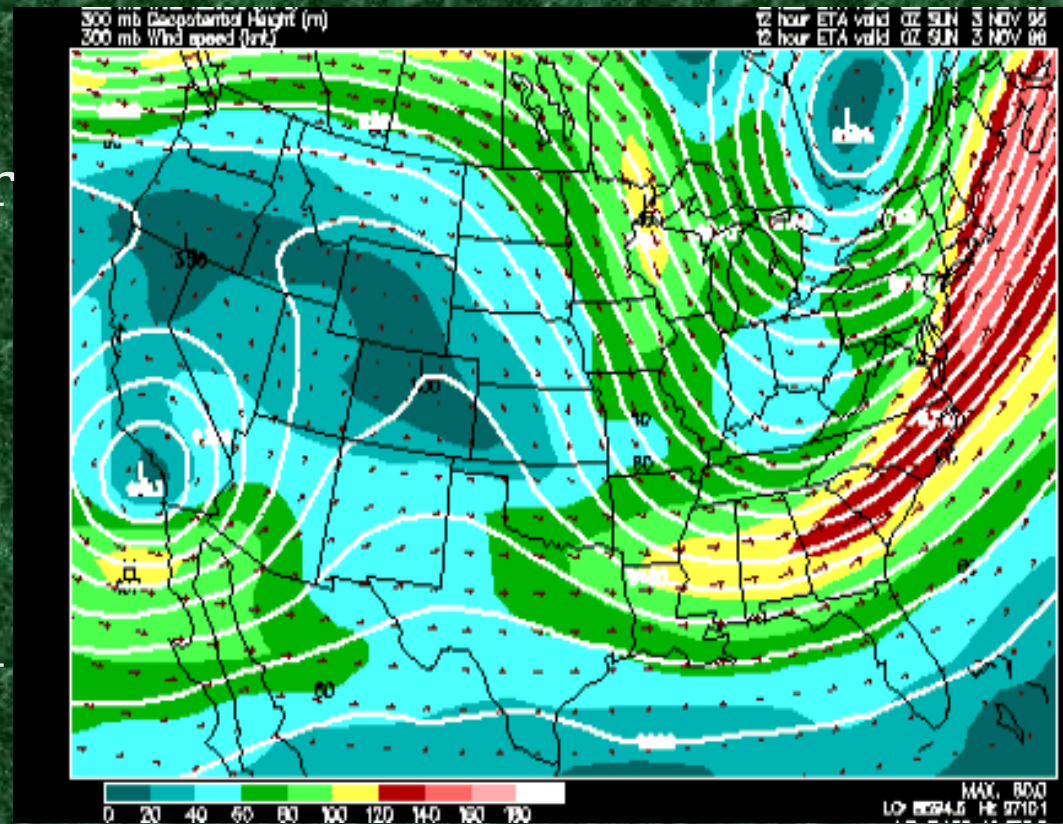
Jet Stream current of rapidly moving air

The jet stream is a current of fast moving air found in the upper levels of the atmosphere. This rapid current is typically thousands of kilometers long, a few hundred kilometers wide, and only a few kilometers thick. Jet streams are usually found somewhere between 10-15 km (6-9 miles) above the earth's surface. The position of this upper-level jet stream is strongly influenced by the position of the strongest SURFACE pressure systems (see the diagram below).



During the winter months, Arctic and tropical air masses create a stronger surface temperature contrast resulting in a strong jet stream. However, during the summer months, when the surface temperature variation is less dramatic, the winds of the jet are weaker.

To the right is an ETA Model forecast panel for 300 mb winds and **geopotential** heights (white contours). The color filled regions indicate wind speed in knots and is color coded according to the legend at the bottom of the image. The shades of blue indicate winds less than 60 knots, while the shades of green, yellow, and red indicate stronger winds.



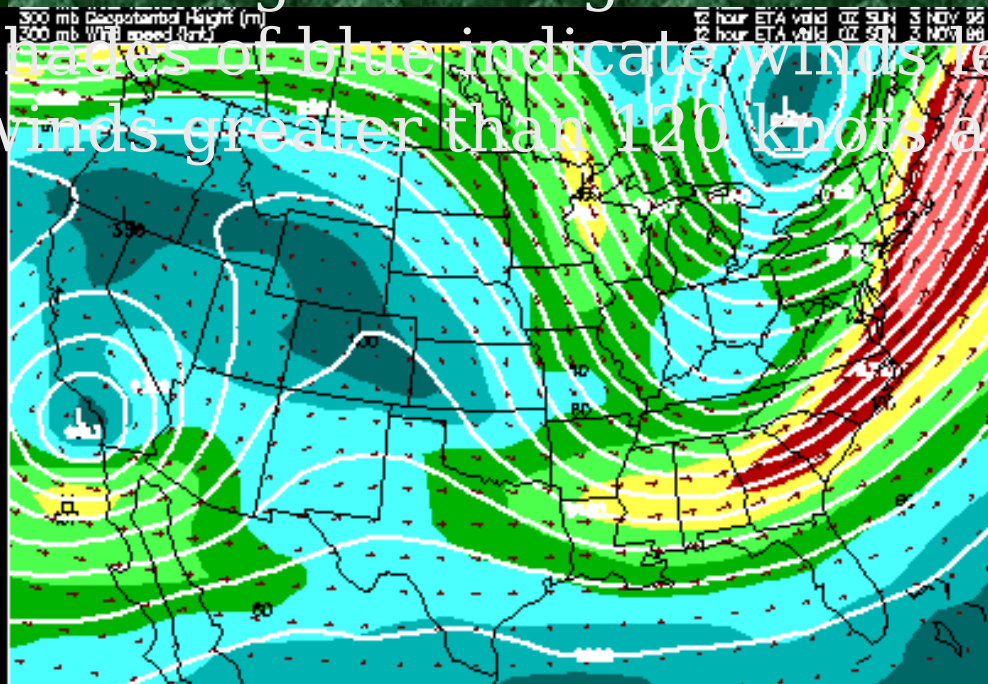
The yellow, green and red ribbon on the image above represents the jet stream, and along the East Coast, the region of strongest winds (shaded in red) is a

Jet Streaks

wind speed maxima within the

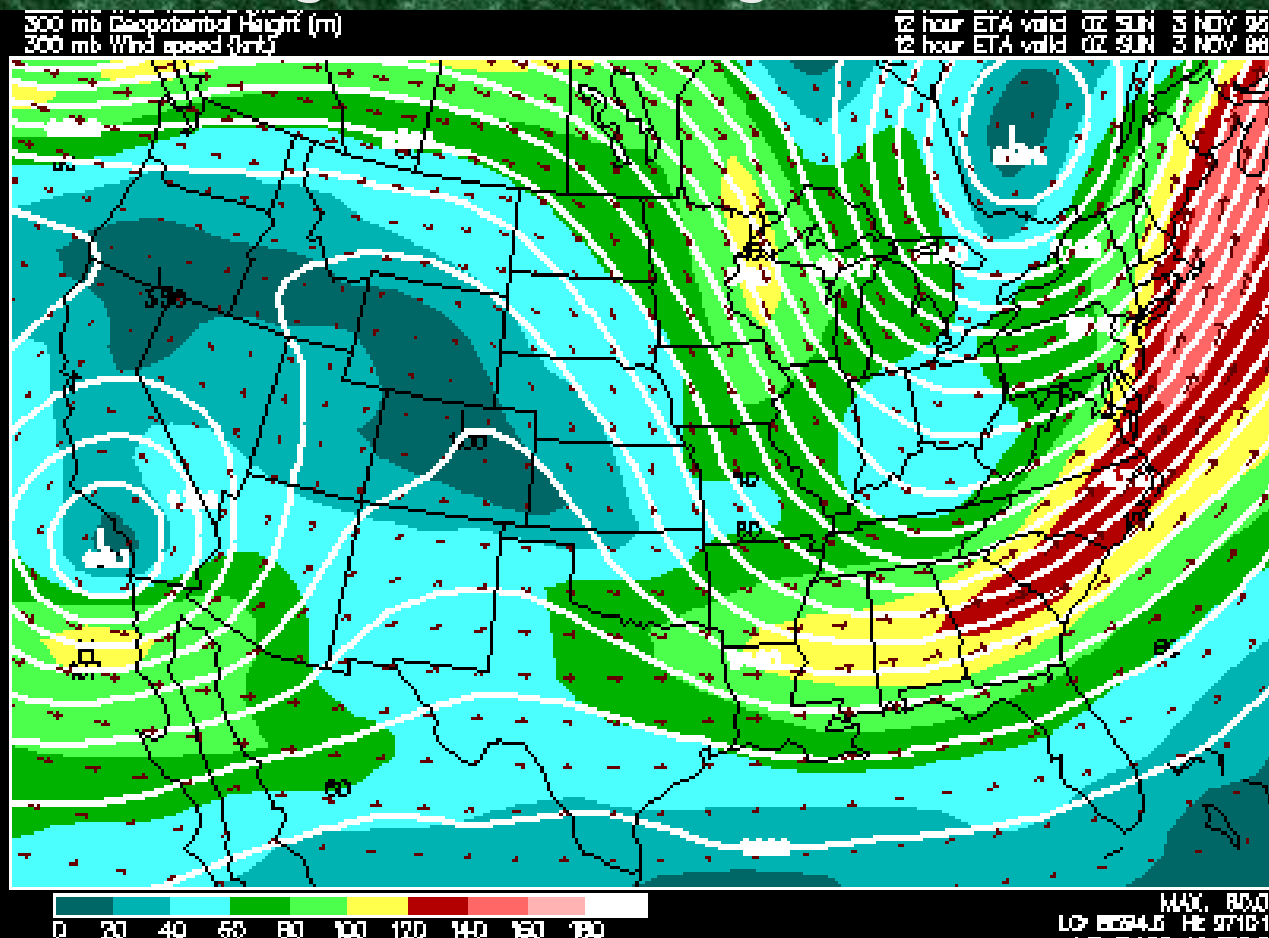
jet stream

Jet streaks are localized regions of very fast winds embedded within the jet stream. Sometimes these local wind maxima reach speeds in excess of 160 knots. Below is an ETA Model forecast panel for 300 mb winds and geopotential heights (white contours). The color filled regions indicate wind speed in knots and is color coded according to the legend at the bottom of the image. The shades of blue indicate winds less than 60 knots, while winds greater than 120 knots are given in shades of red.



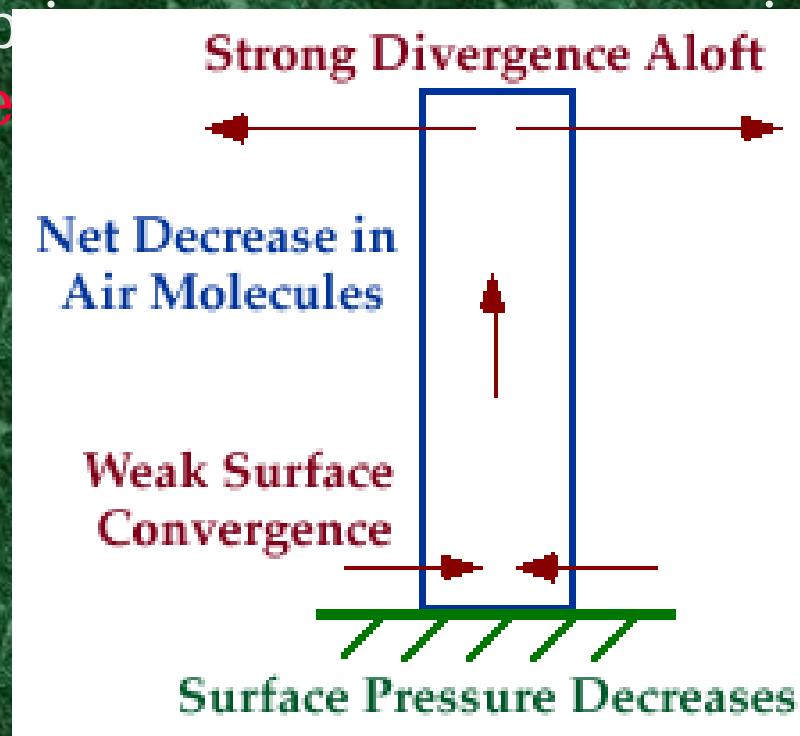
The yellow, green and red ribbon on the image above represent the jet stream, and along the East Coast, the region of strongest winds (shaded in red) is a jet streak.

As air enters a jet streak, it speeds up. When it leaves a jet streak, it slows down. These accelerations and decelerations, coupled with the curvature of the jet stream and strong wind shears, cause air to pile up in some areas (convergence) and spread out (divergence) in others. These regions of divergence and convergence have features.



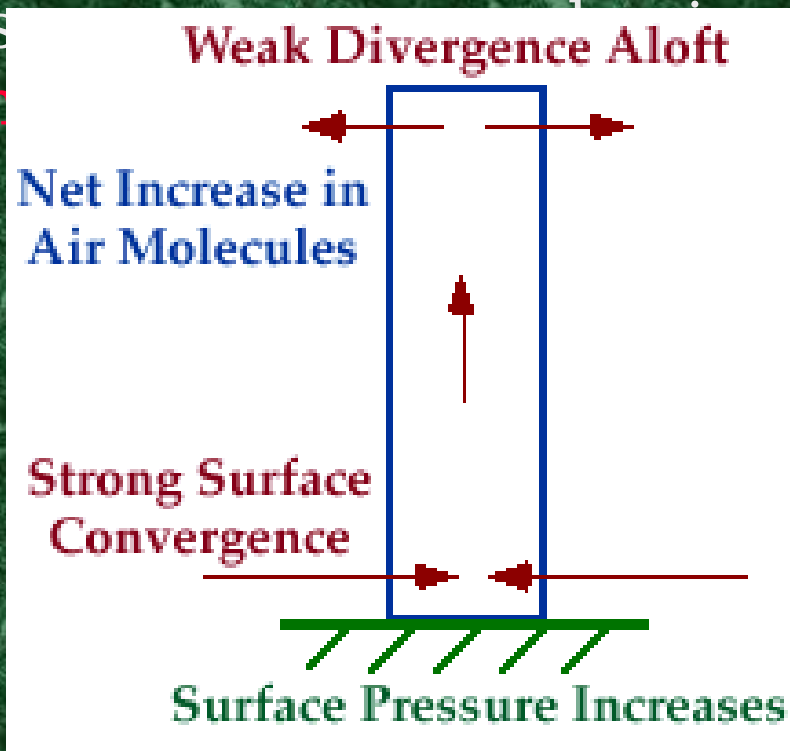
Intensifying Surface Cyclone:

For example, if a region of diverging winds at upper levels is stronger than the converging winds of a surface **low pressure center** below it, the low will deepen (intensify). This is because more air is being removed from the **vertical column of air above the low** than flowing into it, causing the pressure at the surface to decrease. A drop in the **low pressure**



Weakening Surface Cyclone:

In contrast, if a region of diverging winds at upper levels is weaker than the converging winds of a surface **low pressure center** below it, the low begins to fill (weaken). This is because more air is flowing into the **vertical column of air above the low** than flowing out of it, causing the pressure at the surface to increase. An increase in pressure at the surface of the **low pressure center**



Vertical Motion and Jet Streaks

associated with different regions of a jet streak

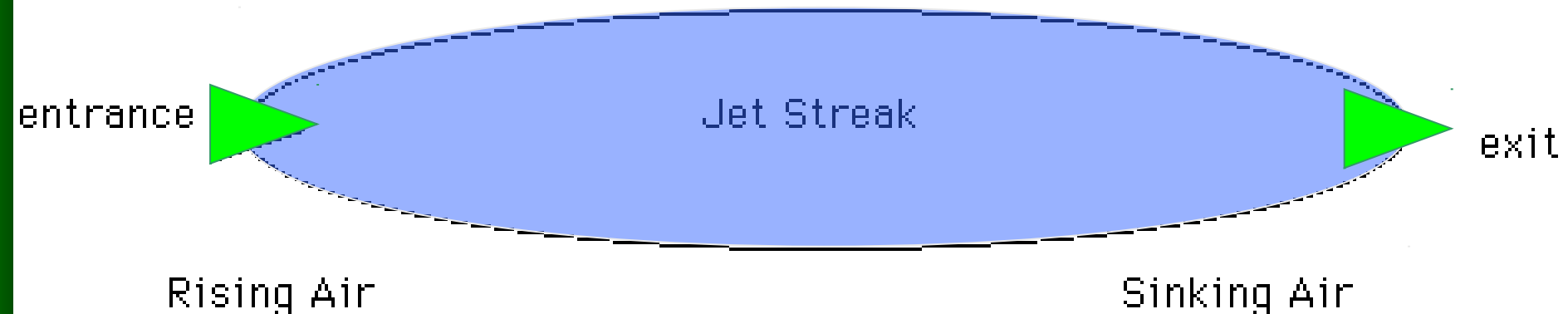
Jet streaks are important as they are indicative of rising motion/falling **pressures** at the surface. The figure below represents an idealized jet streak.

Convergence
Do

Divergence
Don't

Sinking Air

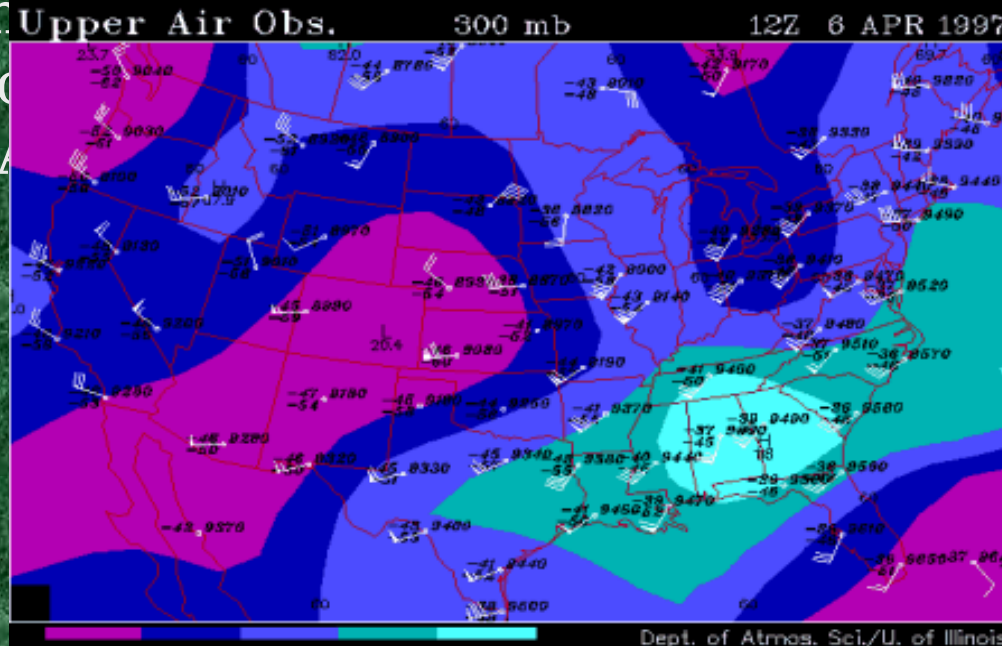
Rising Air



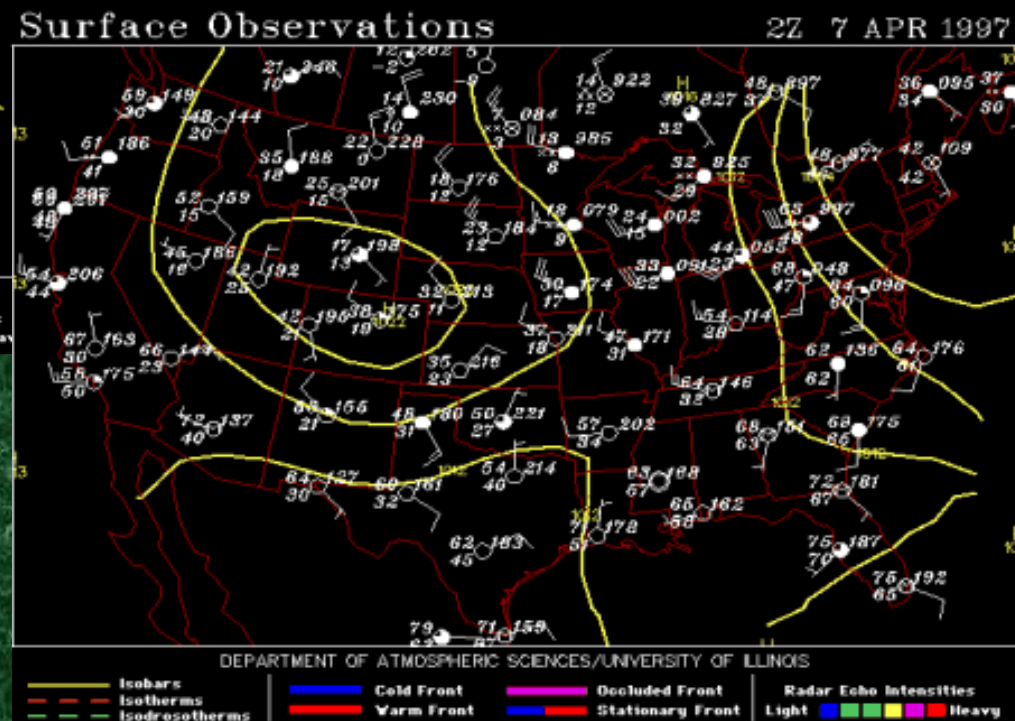
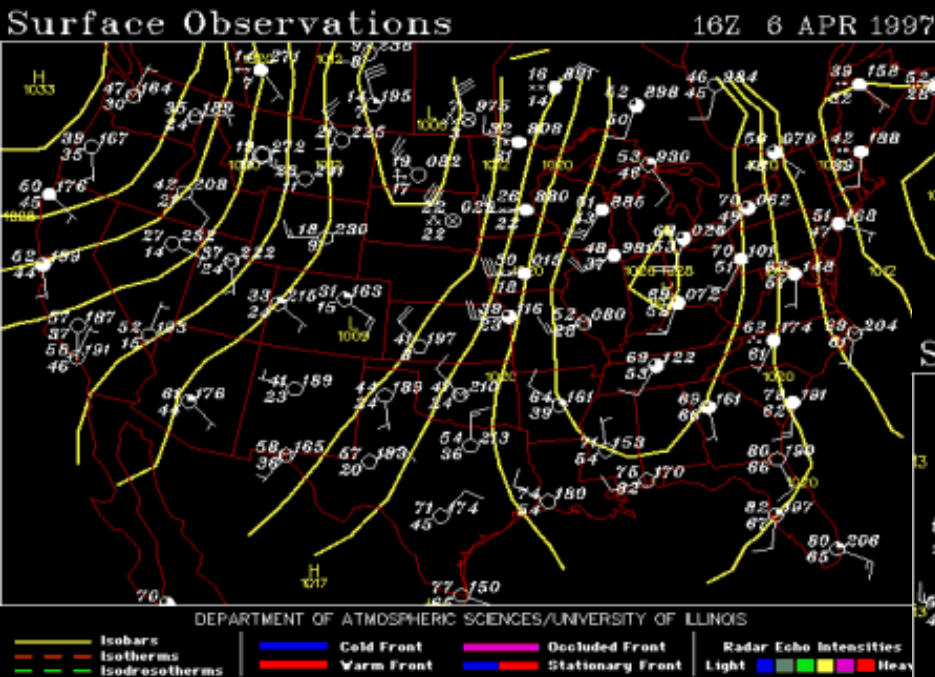
Do
Divergence

Do
Convergence

As air enters from the left, it must be accelerated. The force to do this is supplied by the **Coriolis** force as air flows from the south to the north near the jet entrance, leading to a force to the east (the right). This air motion results in a convergence to the north and a divergence to the south. As a result, air sinks in the northern 'quadrant', and rises in the southern quadrant, leading to pressure changes at the surface. In the jet exit region, the opposite happens, as air flows from north to south to create the force necessary to decelerate the air as it leave the jet streak. The vertical motion resulting from this leads to rising air in the north and sinking air in the south, also leading to pressure changes at the surface.



Notice the jet streak that lies along the southeastern U.S. Now look at the following surface pressure plots.



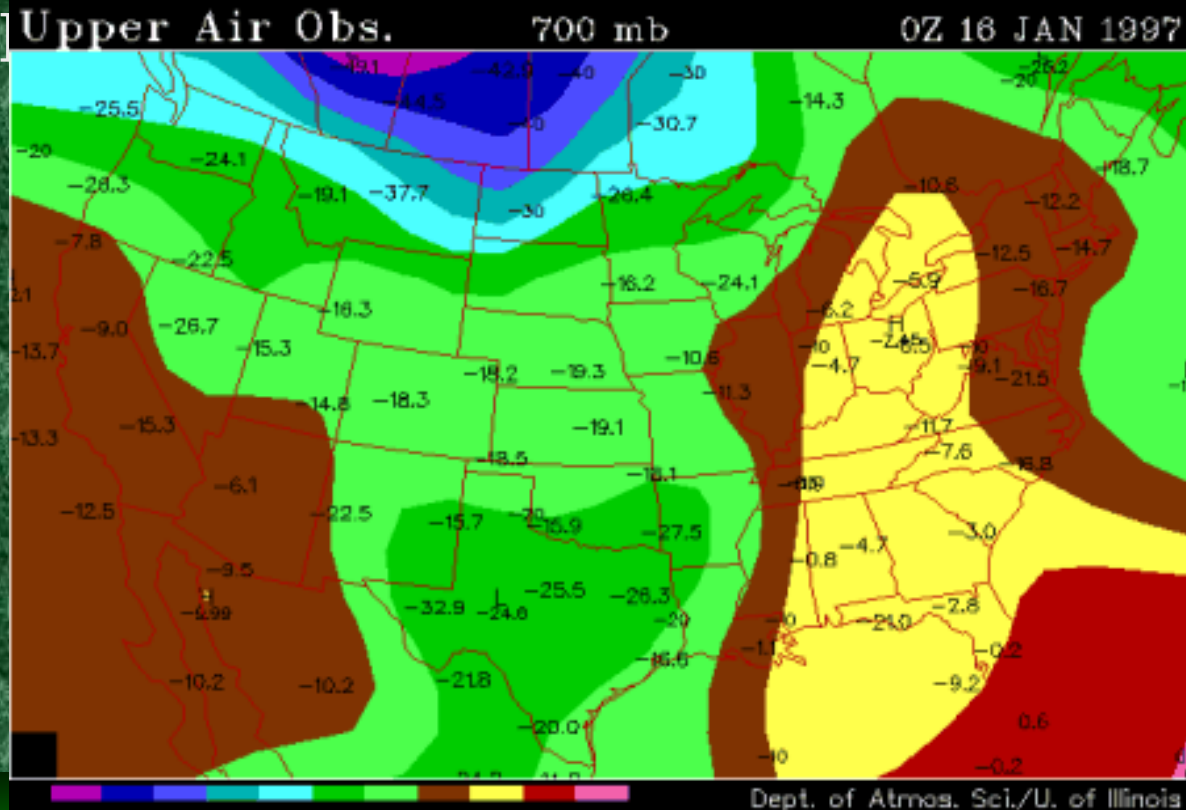
Notice how pressures have risen below the north entrance region, and fallen below the north exit region. This will intensify **cyclones** which tend to be located below the north exit quadrant.

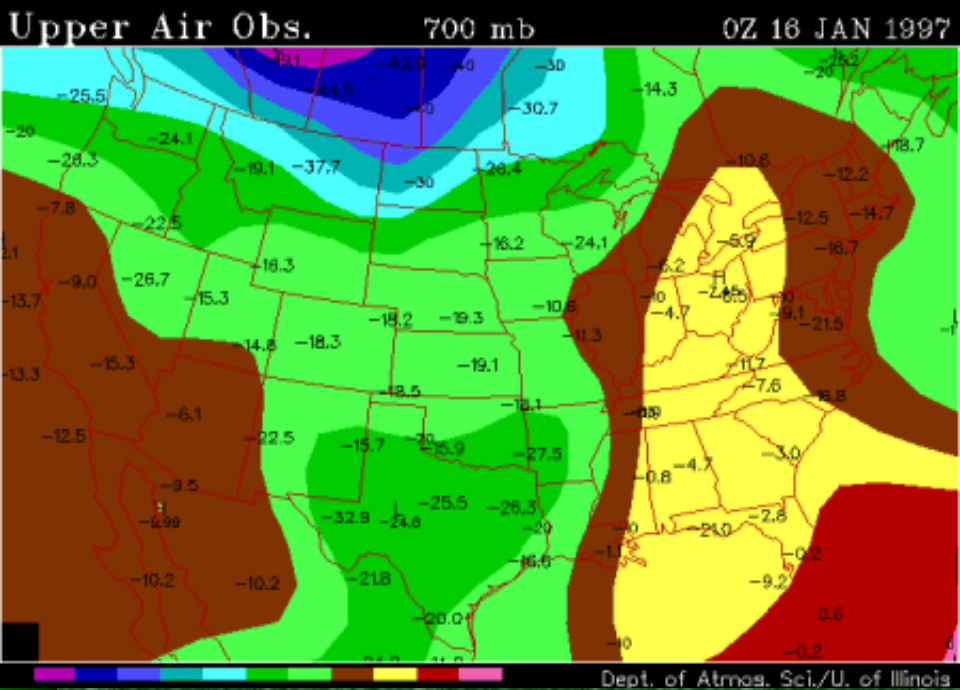
700 mb Dew Point Depression

important indicator for the

presence of clouds

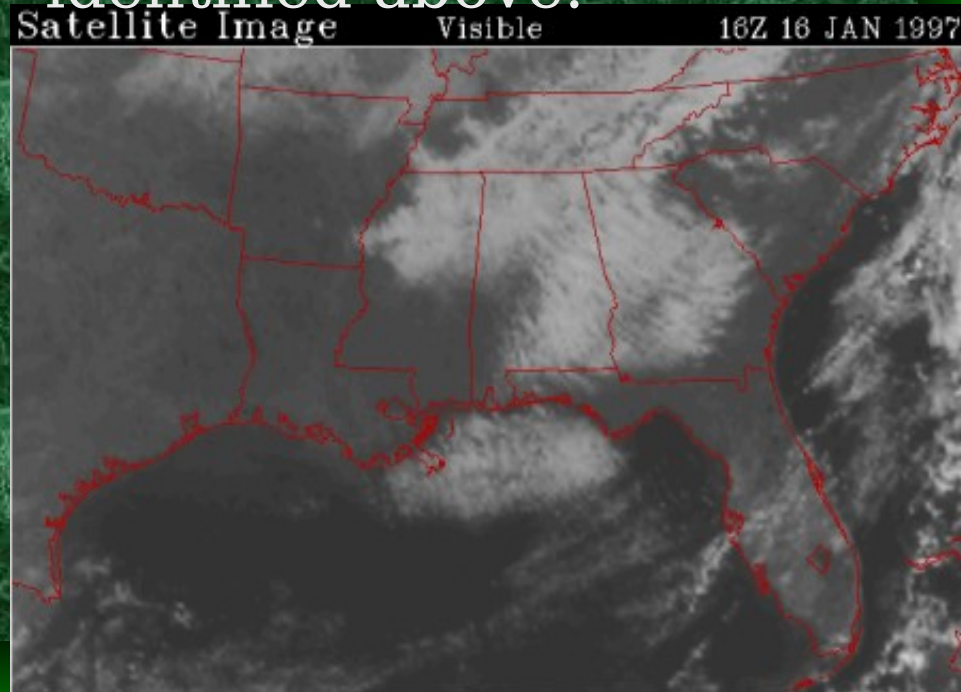
Dew point depression (DD) is useful indicator of how moist the air is and is calculated by taking the difference between the temperature of the air and the dew point temperature. Lower dew point depression values mean that the air is very moist, and an increased likelihood that clouds will develop. In the 700mb map below, the lowest DD values are 1 (in yellow).





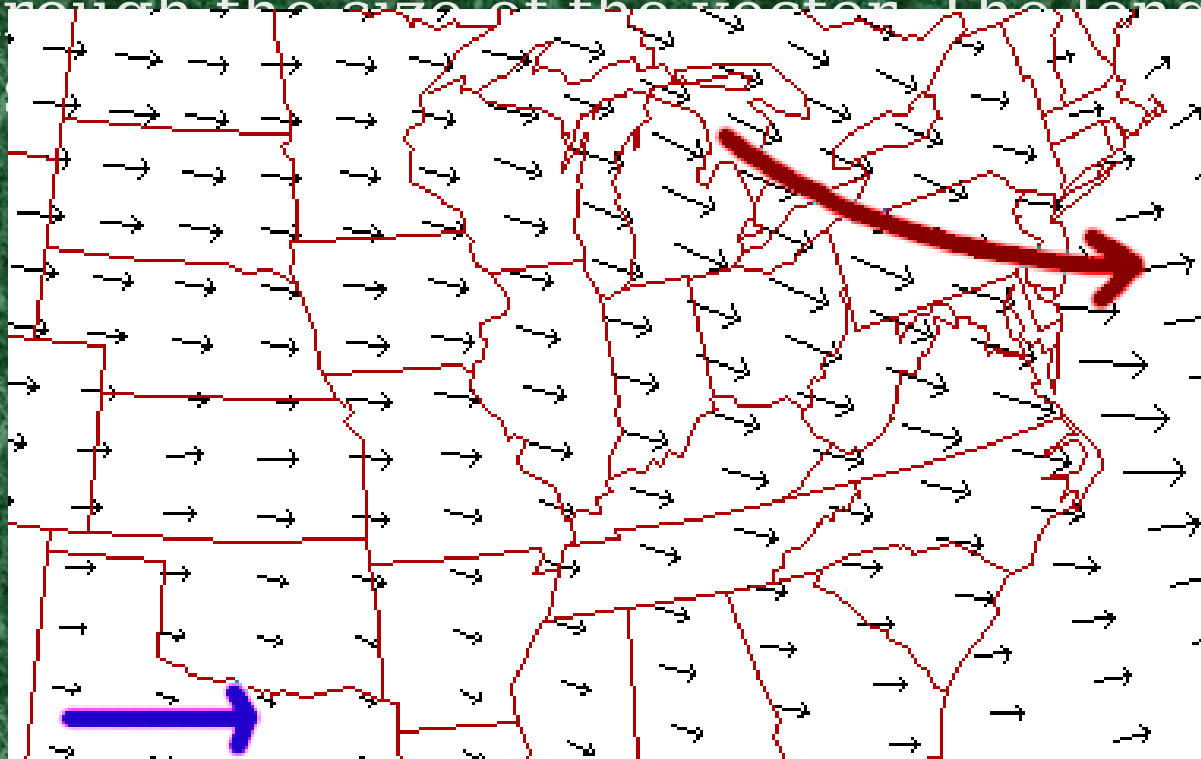
Examination of the **visible satellite image** from the same time reveals an extensive cloud deck present over the Southeast, corresponding very well with the low DD values identified above.

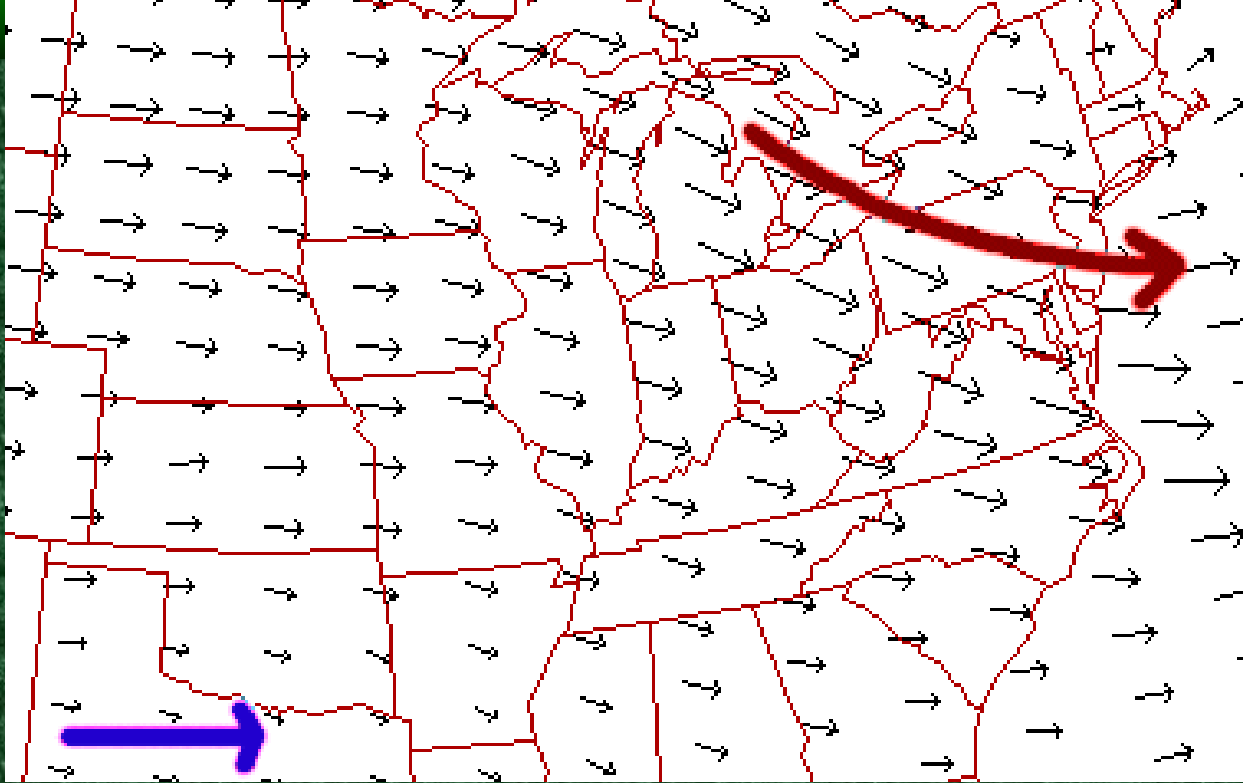
Also observe the lack of clouds over Texas and Louisiana, areas of much greater DD, indicating drier air.



Wind
Vectors
indicate wind direction
and speed

The black arrows plotted on this image are wind vectors. These vectors indicate direction and intensity of the wind. The vectors point in the direction to which the wind is blowing and in this image, winds are primarily blowing from west to east. Intensity of the wind is conveyed through the size of the vector. The longer the arrows, the

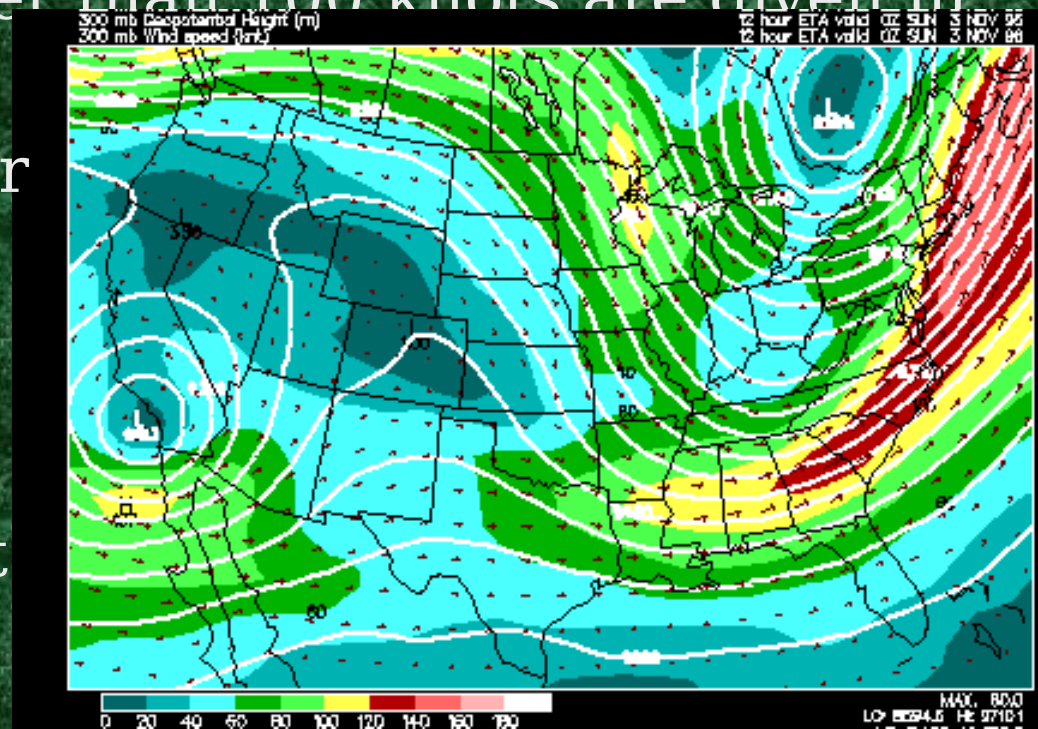




For example, wind vectors in the vicinity of the **red** arrow are longer than those near the **blue** arrow. This means that by the **red** arrow, the winds are stronger, than by the **blue** arrow. Wind vectors are also useful in finding regions of upper level convergence and divergence, which indicate regions of upward and downward motion. Upward motion is typically associated with clouds and precipitation.

Below is an ETA Model forecast panel for 300 mb winds and **geopotential** heights (white contours). The color filled regions indicate wind speed in knots and is color coded according to the legend at the bottom of the image. The shades of blue indicate winds less than 60 knots, while winds greater than 100 knots are given in

shades of red. The wind vectors (red arrows) are much smaller in the blue regions, where the winds are relatively weak, and largest in the red, which is the region of strongest winds. The ribbon of strongest winds (green, yellow and red colors) is called the jet stream with a jet maximum, or jet



Notice how the wind vectors are aligned generally parallel to the geopotential height contours. Once high enough above the earth, the effects of surface friction on wind direction decrease dramatically and